

DATASET description metadata for BCO-DMO

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Lead PI: Robert F. Anderson; U. S. GEOTRACES North Atlantic Transect (GNAZT); Cruises KN199-04, KN204-01A, and KN204-01B; Seawater dissolved and small and large particulate ^{232}Th , ^{230}Th , and ^{231}Pa

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Dataset Name:

GNAZT Th/Pa

Dataset description:

Depth profiles of seawater dissolved and small and large particulate ^{232}Th , ^{230}Th , and ^{231}Pa

Project: Collaborative Research: U. S. GEOTRACES North Atlantic Section: Analysis of ^{230}Th , ^{232}Th and ^{231}Pa

Funding: NSF-OCE 0927064

Cruise or Deployment:

The data contained herein were produced from three cruises aboard the U. S. Research Vessel *Knorr*:

KN199-04 (Lisbon to Cape Verde Islands)

KN204-01A (Woods Hole to Bermuda)

KN204-01B (Bermuda to Cape Verde Islands)

Deployment Synonyms: GEOTRACES North Atlantic Transect; GEOTRACES North Atlantic Zonal Transect; GNAZT; GA03

Location: North Atlantic

Parameter names, definitions and units:

Notes:

Radionuclide concentrations are given as micro-Becquerel (10^{-6} Bq, μ Bq or micro-Bq) per kg seawater for ^{230}Th and ^{231}Pa , and pmol (10^{-12} mol) per kg seawater for ^{232}Th . A Becquerel is the SI unit for radioactivity and is defined as 1 disintegration per second. These units are recommended by the GEOTRACES community.

“Dissolved” (D) here refers to that which passed through a $0.45\ \mu\text{m}$ Acropak™ 500 filter capsule sampled from conventional Niskin bottles. This is true for all dissolved samples except for a select number that came from a towed pumping system designed to collect uncontaminated water at 2-3 m depth, indicated by FISH or Stn-GeoF in the sample bottle type. Surface Fish samples were filtered by a $0.2\ \mu\text{m}$ Osmonics filter capsule, unlike Niskin samples.

The “small particulate” (SP) data refers to the particle size class $0.8\text{--}51\ \mu\text{m}$ and is sometimes also called the “suspended” size fraction. The “large particulate” (LP) data refers to particles greater than $51\ \mu\text{m}$ and sometimes referred to as the “sinking” size fraction. The particulate samples were collected by in-situ pumping over paired $0.8\ \mu\text{m}$ Pall Supor800 polyethersulfone filters behind a $51\ \mu\text{m}$ Sefar polyester mesh prefilter (See Lam et al. 2015 for particulate sampling methodology). Analysis of the paired Supor filters represents a particle size class approximating $0.45\text{--}51\ \mu\text{m}$ (Bishop et al. 2012), while the top filter alone represents $0.8\text{--}51\ \mu\text{m}$ and it is this size class referred to here as the small particulate fraction. We measured a select number of top and bottom filters separately for radionuclides and found that the bottom filters had radionuclide levels that were indistinguishable from clean filter process blanks. Therefore whether or not samples were analyzed as top and bottom paired, or the top filter alone, we infer the small particulate data to represent $0.8\text{--}51\ \mu\text{m}$ particles. The large particulate data is based on analysis of the $51\ \mu\text{m}$ Sefar polyester mesh prefilter. Only a small selection of large particle samples (16) were analyzed at the University of Minnesota.

For dissolved, seawater was weighed directly in the laboratory to determine sample size, taking into account acid added at sea. For particulates, sample size was measured by volume (liters) of seawater pumped by mass flow controllers on the in-situ pumps, but we converted seawater volume to seawater mass using a fixed seawater density of $1.025\ \text{kg/L}$. Concentrations below detection are listed as “bdl”. The abbreviation “nd” refers to no data available.

Parameter names in the form such as “Th_232_D_CONC_BOTTLE” are adopted based on a recommendation from the GEOTRACES community (http://www.egeotraces.org/Parameter_Naming_Conventions.html).

This is compiled data produced by three laboratories with the following associations: Lamont-Doherty Earth Observatory of Columbia University (LDEO), Woods Hole Oceanographic Institution (WHOI) and the University of Minnesota (UMN).

The primary reference for the dissolved data is Hayes et al. (2015, Deep-Sea Research Part II) and for the particulate data, Hayes et al. (2015, Marine Chemistry).

Parameter	Description	Units
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SAMPLE INFO:

cruise_id	Ship's cruise designation	text
cruise_part	Leg of cruise	A or B
station_GEOTRC	Station number, specific to each cruise	integer
LATITUDE	Station latitude	degrees north
LONGITUDE	Station longitude	degrees east
depth_GEOTRC	Sample depth	meters
cast_GEOTRC	Cast number, within a station	integer
event_GEOTRC	GEOTRACES event number	integer
date_collected	Date sample collected	yyyymmdd
time	Time sample collected	hhmm
bottle_GEOTRC	Sample bottle designation, Niskin (NIS), In-situ Pump (PUMP) or towed fish (FISH)	text-integer
sample_GEOTRC	GEOTRACES sample number	integer
lab	Analytical lab (LDEO, UMN or WHOI)	text
date_dissolved_U_separation	Date when dissolved U separated from Th and Pa (used for correcting measured dissolved Th-230 and Pa-231 for ingrowth by decay of dissolved uranium during the time between sample collection and U separation	mm/dd/yyyy

MEASURD PARAMETERS:

Th_232_D_CONC_BOTTLE	Dissolved Th-232 concentration	pmol/kg
Th_232_D_CONC_BOTTLE_ERR	1 sigma error in Dissolved Th-232	pmol/kg
Th_232_D_CONC_BOTTLE_FLAG	Flag for Dissolved Th-232	1 = good, 2 = questionable, 3 = bad or no data
Th_230_D_CONC_BOTTLE	Dissolved Th-230 concentration	μBq/kg
Th_230_D_CONC_BOTTLE_ERR	1 sigma error in Dissolved Th-230	μBq/kg
Th_230_D_CONC_BOTTLE_FLAG	Flag for Dissolved Th-230	1 = good, 2 = questionable, 3 = bad or no data
Pa_231_D_CONC_BOTTLE	Dissolved Pa-231 concentration	μBq/kg
Pa_231_D_CONC_BOTTLE_ERR	1 sigma error in Dissolved Pa-231	μBq/kg
Pa_231_D_CONC_BOTTLE_FLAG	Flag for Dissolved Pa-231	1 = good, 2 = questionable, 3 = bad or no data
Th_232_SP_CONC_PUMP	Small particulate Th-232 concentration	pmol/kg
Th_232_SP_CONC_PUMP_ERR	1 sigma error in small particulate Th-232	pmol/kg
Th_232_SP_CONC_PUMP_FLAG	Flag for small particulate Th-232	1 = good, 2 = questionable, 3 bad or no data
Th_230_SP_CONC_PUMP	Small particulate Th-230 concentration	μBq/kg
Th_230_SP_CONC_PUMP_ERR	1 sigma error in small particulate Th-230	μBq/kg
Th_230_SP_CONC_PUMP_FLAG	Flag for small particulate Th-230	1 = good, 2 = questionable, 3 bad or no data
Pa_231_SP_CONC_PUMP	Small particulate Pa-231 concentration	μBq/kg
Pa_231_SP_CONC_PUMP_ERR	1 sigma error in small particulate Pa-231	μBq/kg
Pa_231_SP_CONC_PUMP_FLAG	Flag for small particulate Pa-231	1 = good, 2 = questionable, 3 bad or no data
Th_232_LP_CONC_PUMP	Large particulate Th-232 concentration	pmol/kg
Th_232_LP_CONC_PUMP_ERR	1 sigma error in large particulate Th-232	pmol/kg
Th_232_LP_CONC_PUMP_FLAG	Flag for large particulate Th-232	1 = good, 2 = questionable, 3 bad or no data

Th_230_LP_CONC_PUMP	Large particulate Th-230 concentration	μBq/kg
Th_230_LP_CONC_PUMP_ERR	1 sigma error in large particulate Th-230	μBq/kg
Th_230_LP_CONC_PUMP_FLAG	Flag for large particulate Th-230	1 = good, 2 = questionable, 3 bad or no data
Pa_231_LP_CONC_PUMP	Large particulate Pa-231 concentration	μBq/kg
Pa_231_LP_CONC_PUMP_ERR	1 sigma error in large particulate Pa-231	μBq/kg
Pa_231_LP_CONC_PUMP_FLAG	Flag for large particulate Pa-231	1 = good, 2 = questionable, 3 bad or no data

DERIVED VARIABLES (unflagged—all are based solely on measured parameters. Units and error designations are the same as for the associated measured parameters. See derived parameter section below for equations used to calculated derived variables):

Th_230_D_XS_CONC_BOTTLE	Dissolved Th-230 concentration corrected for the dissolution of lithogenic minerals, thereby isolating the dissolved Th-230 produced by decay of dissolved uranium
Th_230_D_XS_CONC_BOTTLE_ERR	
Pa_231_D_XS_CONC_BOTTLE	Dissolved Pa-231 concentration corrected for the dissolution of lithogenic minerals (see notes for full explanation)
Pa_231_D_XS_CONC_BOTTLE_ERR	
Th_230_SP_ADS_CONC_PUMP	Small particulate adsorbed Th-230 (see notes for full explanation)
Th_230_SP_ADS_CONC_PUMP_ERR	
Th_230_SP_XS_CONC_PUMP	Small particulate excess Th-230 (see notes for full explanation)
Th_230_SP_XS_CONC_PUMP_ERR	
Pa_231_SP_ADS_CONC_PUMP	Small particulate adsorbed Pa-231 (see notes for full explanation)
Pa_231_SP_ADS_CONC_PUMP_ERR	
Pa_231_SP_XS_CONC_PUMP	Small particulate excess Pa-231 (see notes for full explanation)
Pa_231_SP_XS_CONC_PUMP_ERR	
Th_230_LP_ADS_CONC_PUMP	Large particulate adsorbed Th-230 (see notes for full explanation)
Th_230_LP_ADS_CONC_PUMP_ERR	
Th_230_LP_XS_CONC_PUMP	Large particulate excess Th-230 (see notes for full explanation)
Th_230_LP_XS_CONC_PUMP_ERR	
Pa_231_LP_ADS_CONC_PUMP	Large particulate adsorbed Pa-231 (see notes for full explanation)
Pa_231_LP_ADS_CONC_PUMP_ERR	
Pa_231_LP_XS_CONC_PUMP	Large particulate excess Pa-231 (see notes for full explanation)
Pa_231_LP_XS_CONC_PUMP_ERR	

Notes on DERIVED PARAMETERS:

Th_230_D_XS_CONC_BOTTLE:

The dissolved excess Th-230 concentration refers to the measured dissolved Th-230 corrected for a contribution of Th-230 due to the partial dissolution of uranium-bearing minerals, or lithogenics. Thereby the dissolved excess represents solely the fraction of Th-230 produced in the water by decay of dissolved uranium-234. We estimate the lithogenic Th-230 using measuring dissolved Th-232 and a lithogenic Th-230/Th-232 ratio of 4.0e-6 (atom ratio) as determined by Roy-Barman et al. (2002) and a conversion factor to convert picomoles to micro-Becquerels.

$$\text{Th_230_D_XS_CONC_BOTTLE} = \text{Th_230_D_CONC_BOTTLE} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_D_CONC_BOTTLE}$$

Pa_231_D_XS_CONC_BOTTLE:

The dissolved excess Pa-231 concentration refers to the measured dissolved Pa-231 corrected for a contribution of Pa-231 due to the partial dissolution of uranium-bearing minerals, or lithogenics. Thereby the dissolved excess represents solely the fraction of Pa-231 produced in the water by decay of dissolved uranium-235. We estimate the lithogenic Pa-231 using measuring dissolved Th-232 and a lithogenic Pa-231/Th-232 ratio of 8.8e-8 (atom ratio) which is derived from assuming an average upper continental crustal U/Th ratio (Taylor and McClelland, 1995) and secular equilibrium between Pa-231 and U-235 in the lithogenic material. An additional conversion factor is needed to convert picomoles to micro-Becquerels.

$$\text{Pa_231_D_XS_CONC_BOTTLE} = \text{Pa_231_D_CONC_BOTTLE} - 8.8\text{e-}8 * 4.0370\text{e}5 * \text{Th_232_D_CONC_BOTTLE}$$

Th_230_SP_ADS_CONC_PUMP:

The small particulate adsorbed Th-230 concentration refers to the measured small particulate Th-230 corrected for a contribution of Th-230 locked within mineral lattices. For the particulate Th-230 supported by decay of U within mineral lattices, we use measured small particulate Th-232 and a lithogenic Th-230/Th-232 atomic ratio of 4.0e-6. However, because some fraction of the Th-232 is also adsorbed (and not merely found within mineral lattices), we use the dissolved Th-232/Th-230 ratio to estimate what fraction of the small particulate Th-232 is adsorbed and furthermore to calculate the correction for adsorbed Th-230. Conversion factors are also necessary to convert picomoles to micro-Becquerels. When dissolved data do not exist at the same depth of the particulate samples, we linearly interpolated the dissolved data onto the depths of the particulate samples. See Hayes et al. (2015, Marine Chemistry) for more details.

$$\text{Th_230_SP_ADS_CONC_PUMP} = \frac{[\text{Th_230_SP_CONC_PUMP} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_SP_CONC_PUMP}]}{[1 - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_D_CONC_BOTTLE} / \text{Th_230_D_CONC_BOTTLE}]}$$

Th_230_SP_XS_CONC_PUMP:

The small particulate excess Th-230 refers to the measured small particulate Th-230 corrected for a contribution of Th-230 originating from U-bearing minerals or lithogenic Th-230. Some of this lithogenic Th-230 will be still intact within minerals and some after partial dissolution will have adsorbed to particle surface to contribute in part to the total adsorbed Th-230. Using the measured small particulate Th-232 and a lithogenic Th-230/232 atomic ratio of 4e-6, and not taking into

account the fact that some of the measured particulate Th-232 is adsorbed, corrects for all of the lithogenic Th-230 whether it be adsorbed or within intact minerals. This excess Th-230 is what should be used in scavenging or particle flux studies where it is desired to compare particulate Th-230 concentrations to Th-230 production by decay of uranium dissolved in seawater. An additional conversion factor converts picomoles to micro-Becquerels.

$$\text{Th_230_SP_XS_CONC_PUMP} = \text{Th_230_SP_CONC_PUMP} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_SP_CONC_PUMP}$$

Pa_231_SP_ADS_CONC_PUMP:

The small particulate adsorbed Pa-231 concentration refers to the measured small particulate Pa-231 corrected for a contribution of Pa-231 locked within mineral lattices (as opposed to adsorbed on to particle surfaces). The particulate Pa-231 supported by decay of U within mineral lattices, we use measured small particulate Th-232 and a lithogenic Pa-231/Th-232 atomic ratio of $8.8\text{e-}8$. However, because some fraction of the Th-232 is also adsorbed (and not merely found within mineral lattices), we use the dissolved and particulate Th-232/Th-230 ratio to estimate what fraction of the small particulate Th-232 is adsorbed and furthermore to calculate the correction for adsorbed Pa-231. Conversion factors are also necessary to convert picomoles to micro-Becquerels. When dissolved data do not exist at the same depth of the particulate samples, we linearly interpolate the dissolved data onto the depths of the particulate samples. See Hayes et al. (2015, Marine Chemistry) for more details.

$$\begin{aligned} \text{Pa_231_SP_ADS_CONC_PUMP} = & \text{Pa_231_SP_CONC_PUMP} - \\ & 8.8\text{e-}8 * 4.0370\text{e}5 * (\text{Th_232_SP_CONC_PUMP} - \text{Th_232_D_CONC_BOTTLE} * \\ & [(\text{Th_230_SP_CONC_PUMP} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_SP_CONC_PUMP}) / \\ & (\text{Th_230_D_CONC_BOTTLE} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_D_CONC_BOTTLE})]) \end{aligned}$$

Pa_231_SP_XS_CONC_PUMP:

The small particulate excess Pa-231 refers to the measured small particulate Pa-231 corrected for a contribution of Pa-231 originating from U-bearing minerals or lithogenic Pa-231. Some of this lithogenic Pa-231 will be still intact within minerals and some after partial dissolution will have adsorbed to particle surface to contribute in part to the total adsorbed Pa-231. Using the measured small particulate Th-232 and a lithogenic Pa-231/Th-232 atomic ratio of $8.8\text{e-}8$, and not taking into account the fact that some of the measured particulate Th-232 is adsorbed, corrects for all of the lithogenic Pa-231 whether it be adsorbed or within intact minerals. This excess Pa-231 is what should be used in scavenging or particle flux studies where it is desired to compare particulate Pa-231 concentrations to Pa-231 production by decay of uranium dissolved in seawater. An additional conversion factor converts picomoles to micro-Becquerels.

$$\text{Pa_231_SP_XS_CONC_PUMP} = \text{Pa_231_SP_CONC_PUMP} - 8.8\text{e-}8 * 4.0370\text{e}5 * \text{Th_232_SP_CONC_PUMP}$$

Th_230_LP_ADS_CONC_PUMP:

The large particulate adsorbed Th-230 concentration refers to the measured large particulate Th-230 corrected for a contribution of Th-230 locked within mineral lattices. For the particulate Th-230 supported by decay of U within mineral lattices, we use measured large particulate Th-232 and a lithogenic Th-230/Th-232 atomic ratio of 4.0×10^{-6} . However, because some fraction of the Th-232 is also adsorbed (and not merely found within mineral lattices), we use the dissolved Th-232/Th-230 ratio to estimate what fraction of the large particulate Th-232 is adsorbed and furthermore to calculate the correction for adsorbed Th-230. Conversion factors are also necessary to convert picomoles to micro-Becquerels. When dissolved data do not exist at the same depth of the particulate samples, we linearly interpolated the dissolved data onto the depths of the particulate samples. See Hayes et al. (2015, Marine Chemistry) for more details.

$$\text{Th}_{230_LP_ADS_CONC_PUMP} = \frac{[\text{Th}_{230_LP_CONC_PUMP} - 4.0 \times 10^{-6} * 1.7473 \times 10^5 * \text{Th}_{232_LP_CONC_PUMP}]}{[1 - 4.0 \times 10^{-6} * 1.7473 \times 10^5 * \text{Th}_{232_D_CONC_BOTTLE} / \text{Th}_{230_D_CONC_BOTTLE}]}$$

Th₂₃₀_LP_XS_CONC_PUMP:

The large particulate excess Th-230 refers to the measured large particulate Th-230 corrected for a contribution of Th-230 originating from U-bearing minerals or lithogenic Th-230. Some of this lithogenic Th-230 will be still intact within minerals and some after partial dissolution will have adsorbed to particle surface to contribute in part to the total adsorbed Th-230. Using the measured large particulate Th-232 and a lithogenic Th-230/232 atomic ratio of 4×10^{-6} , and not taking into account the fact that some of the measured particulate Th-232 is adsorbed, corrects for all of the lithogenic Th-230 whether it be adsorbed or within intact minerals. This excess Th-230 is what should be used in scavenging or particle flux studies where it is desired to compare particulate Th-230 concentrations to Th-230 production by decay of uranium dissolved in seawater. An additional conversion factor converts picomoles to micro-Becquerels.

$$\text{Th}_{230_LP_XS_CONC_PUMP} = \text{Th}_{230_LP_CONC_PUMP} - 4.0 \times 10^{-6} * 1.7473 \times 10^5 * \text{Th}_{232_LP_CONC_PUMP}$$

Pa₂₃₁_LP_ADS_CONC_PUMP:

The large particulate adsorbed Pa-231 concentration refers to the measured large particulate Pa-231 corrected for a contribution of Pa-231 locked within mineral lattices (as opposed to adsorbed on to particle surfaces). The particulate Pa-231 supported by decay of U within mineral lattices, we use measured large particulate Th-232 and a lithogenic Pa-231/Th-232 atomic ratio of 8.8×10^{-8} . However, because some fraction of the Th-232 is also adsorbed (and not merely found within mineral lattices), we use the dissolved and particulate Th-232/Th-230 ratio to estimate what fraction of the large particulate Th-232 is adsorbed and furthermore to calculate the correction for adsorbed Pa-231. Conversion factors are also necessary to convert picomoles to micro-Becquerels. When dissolved data do not exist at the same depth of the particulate samples, we linearly interpolate the dissolved data onto the depths of the particulate samples. See Hayes et al. (2015, Marine Chemistry) for more details.

$$\text{Pa}_{231_LP_ADS_CONC_PUMP} = \text{Pa}_{231_LP_CONC_PUMP} -$$

$$8.8\text{e-}8 * 4.0370\text{e}5 * (\text{Th_232_LP_CONC_PUMP} - \text{Th_232_D_CONC_BOTTLE} * [(\text{Th_230_LP_CONC_PUMP} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_LP_CONC_PUMP}) / (\text{Th_230_D_CONC_BOTTLE} - 4.0\text{e-}6 * 1.7473\text{e}5 * \text{Th_232_D_CONC_BOTTLE})])$$

Pa₂₃₁_LP_XS_CONC_PUMP:

The large particulate excess Pa-231 refers to the measured large particulate Pa-231 corrected for a contribution of Pa-231 originating from U-bearing minerals or lithogenic Pa-231. Some of this lithogenic Pa-231 will be still intact within minerals and some after partial dissolution will have adsorbed to particle surface to contribute in part to the total adsorbed Pa-231. Using the measured large particulate Th-232 and a lithogenic Pa-231/Th-232 atomic ratio of 8.8e-8, and not taking into account the fact that some of the measured particulate Th-232 is adsorbed, corrects for all of the lithogenic Pa-231 whether it be adsorbed or within intact minerals. This excess Pa-231 is what should be used in scavenging or particle flux studies where it is desired to compare particulate Pa-231 concentrations to Pa-231 production by decay of uranium dissolved in seawater. An additional conversion factor converts picomoles to micro-Becquerels.

$$\text{Pa_231_LP_XS_CONC_PUMP} = \text{Pa_231_LP_CONC_PUMP} - 8.8\text{e-}8 * 4.0370\text{e}5 * \text{Th_232_LP_CONC_PUMP}$$

Sampling Methodology:

Dissolved data:

Water samples were collected with a Sea-Bird Electronics CTD carousel fitted with 12 30-liter PVC Niskin bottles, maintained and operated by the Ocean Data Facility of Scripps Institution of Oceanography. The carousel was lowered from the ship with steel wire. Niskin bottles were equipped with nylon-coated closure springs and Viton O-rings. After collection seawater was drained with Teflon-lined Tygon™ tubing and filtered through Pall Acropak™ 500 filters on deck (gravity filtration, 0.8/0.45 µm pore size) into Fisher I-Chem series 300 LDPE cubitainers. Approximately 4-5 L was collected per desired depth. Prior to the cruise, the tubing, filters and cubitainers were cleaned by immersion in 1.2 M HCl (Fisher Scientific Trace Metal Grade) for 4-5 days. Once filtered, samples were adjusted to a pH ~2 with ultra-clean 6 M HCl (Fisher Scientific OPTIMA grade), double-bagged, stored in pallet boxes on-deck until the end of the cruise and then at room temperature once shipped to the participating laboratories for analysis.

Particulate data:

Size-fractionated particles were collected using McLane Research in-situ pumps (WTS-LV) that had been modified to accommodate two flowpaths (Lam and Morris Patent pending). The wire-out was used to target depths during deployment, and a self-recording Seabird 19plus CTD deployed at the end of the line and RBR data loggers attached to three of the eight pumps were used to correct for actual depths during pumping.

Filter holders used were 142 mm-diameter “mini-MULVFS” style filter holders with two stages for two size fractions and multiple baffle systems designed to ensure even particle distribution and prevent particle loss (Bishop et al. 2012). One of two filter holder/flowpaths was loaded with a 51µm Sefar polyester mesh prefilter followed by paired 0.8 µm Pall Supor800 polyethersulfone filters. Each cast also had “dipped blank” filters deployed. These were the full filters sets (prefilter followed by paired Supor filters) sandwiched within a 1 µm polyester mesh filter, loaded into perforated

polypropylene containers, attached with plastic cable ties to a pump frame, and deployed. Dipped blank filters were exposed to seawater for the length of the deployment and processed and analyzed as regular samples, and thus functioned as full seawater process blanks. We analyzed half portions of the top and bottom filters from the “dipped” blank from 1 or more depths for 7 stations.

All filters and filter holders were acid leached prior to use according to methods recommended in the GEOTRACES sample and sample-handling Protocols (Geotraces 2010).

Analytical methods for dissolved radionuclides:

LDEO:

In the on-shore laboratory, samples were weighed to determine sample size, taking into account the weight of the cubitainer and of the acid added at sea. Then weighed aliquots of the artificial isotope yield monitors ^{229}Th (20 pg) and ^{233}Pa (0.5 pg) and 15 mg dissolved Fe were added to each sample. After allowing 1 day for spike equilibration, the pH of each sample was raised to 8-8.5 by adding ~10 mL of concentrated NH_4OH (Fisher Scientific OPTIMA grade) which caused iron (oxy)hydroxide precipitates to form. This precipitate was allowed to settle for 1-2 days before the overlaying seawater was siphoned off. The Fe precipitate was transferred to centrifuge tubes for centrifugation and rinsing with Milli-Q H_2O ($>18\text{ M}\Omega$) to remove the major seawater ions. The precipitate was then dissolved in 16 M HNO_3 (Fisher Scientific OPTIMA grade) and transferred to a Teflon beaker for a high-temperature (180-200°C) digestion with HClO_4 and HF (Fisher Scientific OPTIMA grade) on a hotplate in a HEPA-filtered laminar flow hood. After total dissolution of the sample, another precipitation of iron (oxy)hydroxide followed and the precipitate was washed with Mill-Q H_2O , centrifuged, and dissolved in 12 M HCl for a series of anion-exchange chromatography using 6 mL polypropylene columns each containing a 1 mL bed of Bio-rad resin (AG1-X8, 100-200 mesh size) and a 45 μm porous polyethylene frit (Anderson et al. 2012). The final column elutions were dried down at 180°C in the presence of 2 drops of HClO_4 and taken up in approximately 1 mL of 0.16 M HNO_3 /0.026 M HF for mass spectrometric analysis.

Concentrations of ^{232}Th , ^{230}Th and ^{231}Pa were calculated by isotope dilution using nuclide ratios determined on a VG Elemental AXIOM Single Collector Magnetic Sector ICP-MS with a Resolving Power of ~400 to ensure the highest sensitivity. All measurements were done using a peak jumping routine in ion counting mode. A solution of SRM129, a natural U standard, was run to determine the mass bias correction (assuming that the mass fractionation for Th and Pa are the same as for U). Each sample measurement was bracketed by measurement of an aliquot of the run solution, used to correct for the instrument background count rates on the masses measured. To correct for potential tailing of ^{232}Th into the minor Th and Pa isotopes, beam intensities were measured at the half masses above and below each mass for ^{230}Th , ^{231}Pa , and ^{233}Pa . Tailing under each minor isotope was estimated as the log mean intensity of the half masses on either side of each minor isotope.

Water samples were analyzed in batches of 10-12. Procedural blanks were determined by processing 4-5 L of Milli-Q water in an acid-cleaned cubitainer acidified to pH ~2 with 6 M HCl as a sample in each batch. An aliquot of an intercalibrated working standard solution of ^{232}Th , ^{230}Th and ^{231}Pa , SW STD 2010-1 referred to by Anderson et al. (2012), was added to a separate cubitainer with 5 L of Milli-Q water (acidified to pH 2) and also processed like a sample in each batch. Total procedural blanks for ^{232}Th , ^{230}Th , and ^{231}Pa ranged from 4.2-20.9 pg, 0.8-2.2 fg, and 0-0.85 fg, respectively.

UMN:

In the on-shore laboratory, samples were weighed to determine sample size, taking into account the weight of the cubitainer and of the acid added at sea. Then weighed aliquots of the artificial isotope yield monitors ^{229}Th and ^{233}Pa and several milligrams dissolved Fe were added to each sample. After allowing 3 day for spike equilibration (at a temperature of about 40°C , the pH of each sample was raised to 8-8.5 by adding concentrated NH_4OH which caused iron (oxy)hydroxide precipitates to form. This precipitate was allowed to settle for 1-2 days before the overlaying seawater was siphoned off. The Fe precipitate was transferred to centrifuge tubes for centrifugation and rinsing with deionized H_2O ($>18\text{ M}\Omega$) to remove the major seawater ions. The precipitate was then dissolved in 14 M HNO_3 and transferred to a Teflon beaker. It was then dried down and taken up in 7N HNO_3 for anion-exchange chromatography using AG1-X8, 100-200 mesh size resin and polyethylene frit. Initial separation was done on Teflon columns with a 0.75 ml column volume (CV). The sample was loaded in one CV of 7N HNO_3 , followed by 1.5 CV of 7N HNO_3 , 3 CV of 8N HCl (collect Th fraction), and 3 CV of 8N HCl combined with 0.015N HF (collect Pa fraction). The Pa and Th fractions were then dried down and taken up in 7N HNO_3 . They were each passed through second and third columns (each with 0.5 ml column volumes) using similar elution schemes. The final Pa and Th fractions were then dried down and dissolved in weak nitric acid for analysis on the mass spectrometer.

Concentrations of ^{232}Th , ^{230}Th and ^{231}Pa were calculated by isotope dilution using nuclide ratios determined on a Thermo-Finnigan Neptune mass spectrometer. All measurements were done using a peak jumping routine in ion counting mode on the discreet dynode multiplier behind the retarding potential quadrupole. A solution of ^{233}U - ^{236}U tracer was run to determine the mass bias correction (assuming that the mass fractionation for Th and Pa are the same as for U). Each sample measurement was bracketed by measurement of an aliquot of the run solution used to correct for the instrument background count rates on the masses measured.

Procedural blanks for chemical and mass spectrometric analyses at Minnesota are about 700 fg for ^{232}Th , 15 ag for ^{230}Th , and 20 ag for ^{231}Pa .

Further details on Pa and Th analysis at the U. Minnesota laboratory are given in Shen et al. (2002, 2003, 2012) and Cheng et al. (2000).

WHOI:

In the on-shore laboratory, the analytical procedures followed the protocols described in Auro et al. (2012). Briefly, samples were weighed to determine sample size, taking into account the weight of the cubitainer. Then weighed aliquots of the artificial isotope yield monitors ^{229}Th and ^{233}Pa and ~100 mg dissolved Fe were added to each sample. After allowing 1 day for spike equilibration, the pH of each sample was raised to 7.5-8 by adding ~10 mL of concentrated NH_4OH (Fisher Scientific OPTIMA grade) which caused iron (oxy)hydroxide precipitates to form. This precipitate was allowed to settle for 5-7 days before the overlaying water was siphoned off. The Fe precipitate and remaining water was transferred to polypropylene centrifuge tubes for centrifugation and rinsing with pH 8 Milli-Q H_2O ($>18\text{ M}\Omega$) to remove the major seawater ions. The precipitate was then transferred into Teflon centrifuge tubes, re-rinsed and dissolved in 12 M HCl (Fisher Scientific OPTIMA grade) for a series of anion-exchange chromatography using 10 mL polypropylene columns each containing a 0.5 mL bed of Eichrom Technologies pre-filter resin in addition to 5 mL Anion Exchange Resin (1x8,

100-200 mesh; Eichrom Technologies) for Th elution and 1.5 mL Anion Exchange Resin for Pa. A ^{236}U tracer was added to the Th fraction to assist in monitoring signal intensity changes during mass spectrometry. The final column elutions were dried down at 150°C in the presence of 1 mL of 16 M HNO_3 (Fisher Scientific OPTIMA grade) and taken up in 1 mL of 0.8 M HNO_3 /0.13 M HF for mass spectrometric analysis.

Concentrations of ^{232}Th , ^{230}Th and ^{231}Pa were calculated by isotope dilution using nuclide ratios determined on a Neptune Multi Collector ICP-MS (Auro et al. 2012). Thorium isotopes were measured using a peak jumping routine with ^{229}Th and ^{230}Th analyzed on the central Secondary Electron Multiplier (SEM) and ^{232}Th and ^{236}U measured concurrently on Faraday collectors. Mass bias and ion counter yields were corrected for using an in-house thorium standard. Peak tails were calculated from the ^{232}Th intensity using the pre-determined size of the tail at 2 and 3 amu for ^{230}Th and ^{229}Th respectively. Accuracy was assessed using a secondary consistency standard made at WHOI. ^{231}Pa and ^{233}Pa were analyzed simultaneously on ion counting channels. A solution of CRM145, a natural U standard, was run to determine the mass bias (assuming that the mass fractionation for Th and Pa are the same as for U) and ion counter yields. Sample measurements were bracketed by measurement of an aliquot of the run solution, used to correct for the instrument background count rates on the masses measured. To correct for potential tailing onto the Pa isotopes, beam intensities were measured at the half masses.

Water samples were analyzed in batches of 10-12. Procedural blanks were determined by processing 4-5 L of Milli-Q water in an acid-cleaned cubitainer acidified to pH ~2 with 12 M HCl as a sample in each batch. One or two aliquots of an intercalibrated working standard solution of ^{232}Th , ^{230}Th and ^{231}Pa , SW STD 2010-1 referred to by Anderson et al. (2012), was added to a separate cubitainer with 5 L of Milli-Q water (acidified to pH 2) and also processed like a sample with each batch. Total procedural blanks for ^{232}Th , ^{230}Th , and ^{231}Pa (with the exception of one batch described below) ranged from 3-15 pg, 0.1- 0.9 fg, and 0.1-1.1 fg respectively. One batch had an anomalously high ^{232}Th blank of >2500pg (with ^{230}Th and ^{231}Pa of 4.4 and 0.3 fg respectively) which was found to be due to a batch of contaminated acid. The ^{232}Th concentrations for that batch are not reported.

Further details on sampling and analysis are given by Anderson et al. (2012).

Analytical Methods for particulate radionuclides:

The Supor filters were subsampled in an on-shore laboratory at the Woods Hole Oceanographic Institution and shipped to the participating labs for Pa/Th analysis. Twenty-five to 50% of the paired Supor filters, representing 55-350 L of seawater, were used for Pa/Th analysis. Analyses were similar but differed slightly for the Lamont-Doherty Earth Observatory, WHOI and the University of Minnesota. Details of each groups methodologies can be found in reports by Anderson et al. (2012), Auro et al. (2012) and Shen et al. (2002, 2003, 2012), respectively. Below we give a typical procedure used at LDEO for illustrative purposes.

LDEO procedures:

Filters were folded into 60 mL Teflon jars and weighed aliquots of the artificial isotope yield monitors ^{229}Th (1 pg) and ^{233}Pa (0.3-0.4 pg) and 7-8 mg dissolved Fe were added to each sample. Filters were first heated in ~5 mL 8 N HNO_3 for 1-2 hours at 150°C , then 4-5 mL HClO_4 was added and heat was increased to 200°C until dense white fumes appeared for ~10-20 min. The heat was then

reduced to 180°C and the samples were covered with a Teflon watch cover. After 1-4 hrs, oxidation of the Supor material accelerated, sometimes producing a foam. A foamed sample would be allowed to cool, and re-heated after the beaker walls and watch cover were washed with small amounts of HNO₃ or Milli-Q water. When the Supor material was largely broken down, the watch covers were removed and HF was added in 2 aliquots of ~10-15 drops in between reheating until attaining dense HClO₄ fumes for at least 10 min.

After total dissolution of the sample, the sample-HClO₄ residue was taken up in dilute HCl, and transferred to 50 mL centrifuge tubes with water rinses. Ten to 20 drops of NH₄OH were added to raise pH to 8-8.5 when iron (oxy)hydroxide precipitated. This precipitate was then centrifuged, decanted, washed with Milli-Q H₂O, centrifuged, and dissolved in 12 M HCl for a series of anion-exchange chromatography using 6 mL polypropylene columns each containing a 1 mL bed of Bio-rad resin (AG1-X8, 100-200 mesh size) and a 45 µm porous polyethylene frit (Anderson et al. 2012). The final column elutions were dried down at 180°C in the presence of 2 drops of HClO₄ and taken up in approximately 1 mL of 0.16 M HNO₃/0.026 M HF for mass spectrometric analysis. All acids and bases used were Fisher Chemical OPTIMA grade.

Concentrations of ²³²Th, ²³⁰Th and ²³¹Pa were calculated by isotope dilution using nuclide ratios determined on a Thermo Scientific Element XR Inductively-couple plasma mass spectrometer (ICP-MS) in low resolution. All measurements were done using a peak jumping routine in ion counting mode. A solution of SRM129, a natural U standard, was run to determine the mass bias correction (assuming that the mass fractionation for Th and Pa are the same as for U). Each sample measurement was bracketed by measurement of an aliquot of the run solution, used to correct for the instrument background count rates on the masses measured. To correct for potential tailing of ²³²Th into the minor Th and Pa isotopes, beam intensities were measured at the half masses above and below each mass for ²³⁰Th, ²³¹Pa, and ²³³Pa. Tailing under each minor isotope was estimated as the log mean intensity of the half masses on either side of each minor isotope.

In addition to laboratory procedural blanks (reagents/labware blanks) and periodic measurements of an intercalibrated working standard solution of ²³²Th, ²³⁰Th and ²³¹Pa, SW STD 2010-1 referred to by Anderson et al. (2012), the participating labs also analyzed a number (n = 23) of “dipped blank” filters, mentioned above, to determine the total blank, associated with the sample collection and handling in addition to the laboratory procedure.

For better statistics, we pooled all procedural blank corrected “dipped” blanks (n = 23) to determine filter blank corrections. “Dipped” filter blanks for ²³²Th, ²³⁰Th, and ²³¹Pa were from 156 ± 57 pg, 5.8 ± 2.0 fg, and 0.12 ± 0.04 fg, respectively. Total blanks were < 10% of the measured isotope amounts, except shallower than 200 m water depth, where blanks could be on the order of 50% of the measured ²³⁰Th and ²³¹Pa.

We define the limit of detection as 3 times the standard deviation in the measured “dipped” blanks (170 pg ²³²Th, 6.0 fg ²³⁰Th, and 0.13 fg ²³¹Pa). There were 5 samples for which ²³¹Pa was considered below detection, and all other samples were above the cited limits.

Further details on analysis of seawater particulate radionuclides are given by Anderson et al. (2012).

UMN Procedures:

Filters were folded into 30 mL Teflon beaker and weighed aliquots of the artificial isotope yield monitors ^{229}Th and ^{233}Pa were added. Filters were first completely submerged in 7N HNO_3 acid combined with 10 drops HF, tightly covered with a Teflon threaded cap and heated for 10 hours at 200°F so that the particulate sample was dissolved/leached under pressure. The leach solution was then transferred to a second acid-cleaned Teflon beaker separate from the residual filter. 5 drops of HClO_4 were then added to the leach solution in the second beaker. The original beaker walls and caps were washed with small amounts of weak HNO_3 and the resulting solution added to the second beaker. The solution was then dried down and was taken up in 2N HCl , and transferred to 15ml centrifuge tubes along with a 2N HCl rinse. One drop of dissolved Fe and six to nine drops of NH_4OH were added to raise pH to 8-8.5 at which time iron (oxy)hydroxide precipitated. This precipitate was then centrifuged, decanted, washed with deionized H_2O ($>18\text{ M}\Omega$), centrifuged, and dissolved in 14M HNO_3 and transferred to a Teflon beaker. It was then dried down and taken up in 7N HNO_3 for anion-exchange chromatography using AG1-X8, 100-200 mesh resin and a polyethylene frit. Initial separation was done on Teflon columns (internal diameter $\sim 0.35\text{cm}$) with a $\sim 0.55\text{ ml}$ column volume (CV). The sample was loaded in one CV of 7N HNO_3 , followed by 1.5 CV of 7N HNO_3 , 3 CV of 8N HCl (collect Th fraction), and 3 CV of 8N HCl combined with 0.015N HF (collect Pa fraction). The Pa and Th fractions were then dried down in the presence of 2 drops of HClO_4 and taken up in 7N HNO_3 . They were each passed through second and third columns (each with $\sim 0.55\text{ ml}$ column volumes) using similar elution schemes. The final Pa and Th fractions were then dried down in the presence of 2 drops of HClO_4 and dissolved in weak nitric acid for analysis on the mass spectrometer.

Concentrations of ^{232}Th , ^{230}Th and ^{231}Pa were calculated by isotope dilution using nuclide ratios determined on a Thermo-Finnigan Neptune mass spectrometer. All measurements were done using a peak jumping routine in ion counting mode on the discreet dynode multiplier behind the retarding potential quadrupole. A solution of ^{233}U - ^{236}U tracer was run to determine the mass bias correction (assuming that the mass fractionation for Th and Pa are the same as for U). Each sample measurement was bracketed by measurement of an aliquot of a wash solution, used to correct for the instrument background count rates on the masses measured.

Particulate samples were analyzed in batches of 37 to 39. An aliquot of an intercalibrated working standard solution of ^{232}Th , ^{230}Th and ^{231}Pa , SW STD 2010-1, was added to a separate acid-cleaned Teflon beaker along with weighed aliquots of ^{229}Th spike and ^{233}Pa spike. Spike and Standard were equilibrated for 3 days. The solution was then dried down and taken up in 7N HNO_3 for anion-exchange chromatography using AG1-X8, 100-200 mesh resin and a polyethylene frit, and processed like a sample. In addition to laboratory procedural blanks (reagents/labware blanks), a number of “dipped blank” filters were also processed like samples, to determine the total blank, associated with the sample collection and handling, in addition to the laboratory procedure.

WHOI procedures (Microwave-assisted acid digestion):

Sample digestions were accomplished by microwave-assisted acid digestion performed in a Multiwave 3000 (Anton Paar GmbH, Graz, Austria) instrument. Each sample (1/2 filter) were cut into two quarter filters, and put into two pre-cleaned microwave PFA vessels with 12 mL of Aqua Regia ($\text{HCl} : \text{HNO}_3 = 3:1$) and 0.3 mL of HF for the sample digestion. The microwave procedure consisted of three main steps: (1) a 20 min power increasing ramp; (2) the power was electronically adjusted to maintain the internal temperature (600W) at 200°C for 90 min, and (3) a 20 min power

decreasing ramp. During the session, the temperature ranged from 190°C and 230°C from one vessel to the other.

The two digested quarter filters were combined into one full sample (1/2 filter). Each sample was transferred to 50mL teflon centrifuge tubes with milliQ water rinses. Purified Fe was added to each sample, which was then spiked with ^{229}Th , ^{233}Pa and ^{236}U (Auro et al 2012), and allowed to equilibrate for at least one day. Ten to 20 drops of NH_4OH were added to raise pH to 7.5-8.0 when iron (oxy)hydroxide precipitated. This precipitate was centrifuged, the supernatant decanted, and the remaining precipitate rinsed with pH 8 Milli-Q H_2O , centrifuged, and dissolved in concentrated HCl . The sample was purified and separated using a series of anion-exchange chromatography steps (Auro et al 2012) using 7 mL polypropylene columns each containing Eichrom resin (AG1-X8, 200-400 mesh) and Eichrom pre-filter resin (100-150 μm). The final column elutions were dried down at 180°C and re-dissolved in one drop of concentrated HNO_3 . Dried down and taken up in approximately 1 mL of 5% HNO_3 +0.13N HF and 1 drop of concentrated HNO_3 for mass spectrometric analysis. All acids and bases used were Fisher Chemical OPTIMA grade.

Concentrations of ^{232}Th , ^{230}Th and ^{231}Pa were calculated by isotope dilution (^{229}Th - ^{233}Pa spikes) using isotopic ratios determined on a Thermo Scientific Neptune Multi-Collector Inductively-couple plasma mass spectrometer (MC-ICP-MS) coupled to Cetac Aridus I in Low-Resolution mode (Auro et al 2012). For the Th isotopes, measurements on ^{229}Th and ^{230}Th were done using a “peak-hopping” method where ^{229}Th and ^{230}Th were each analyzed on the central SEM. The ^{232}Th beam was analyzed in both steps on Faraday Cup, allowing direct determinations of $^{232}\text{Th}/^{230}\text{Th}$ and $^{232}\text{Th}/^{229}\text{Th}$. Mass bias correction was assessed using the uranium standard of CRM-145 (assuming that the mass fractionation for Th and Pa are the same as for U), and the accuracy of the method was evaluated by two in-house standards (ThSGS and ThB, Robinson et al., 2005; Auro et al., 2012). For the Pa isotopes, ^{231}Pa and ^{233}Pa were analyzed on the Multi-Ion Counts (MICs) simultaneously. The yield for each ion counters was checked using the mass bias corrected $^{234}\text{U}/^{238}\text{U}$ ratio of CRM-145 uranium standard. To correct for potential tailing of ^{232}Th into the minor Th and Pa isotopes, beam intensities were measured at the half masses above and below each mass for ^{230}Th , ^{231}Pa , and ^{233}Pa . Tailing under each minor isotope was estimated as the log mean intensity of the half masses on either side of each minor isotope.

Data Processing:

The reported errors for radionuclide concentrations represent the propagation of one sigma errors based on the standard isotope ratios collected by ICP-MS, estimated error in the ^{229}Th or ^{233}Pa spike concentration, and the blank correction of the individual isotopes for each sample batch.

Analysis of all samples was completed over the course of several years. A correction was made to account for the ingrowth of ^{230}Th and ^{231}Pa due to the decay of the natural ^{234}U and ^{235}U preserved in the acidified samples during the period of time between sample collection and U-Th/Pa separation during anion exchange chromatography. Thus, the reported ^{230}Th and ^{231}Pa concentrations have been corrected to represent their concentrations at the time of sampling. U concentrations in the samples were estimated using the bottle salinity (S) measured from the CTD and the U-Salinity relationship in seawater (Owens et al. 2011), $[\text{U}] = (0.100 * \text{S} - 0.326) \text{ ng U (g seawater)}^{-1}$. We used seawater U-isotopic compositions of $^{234}\text{U}/^{238}\text{U} = 1.1468$ activity ratio (Andersen et al., 2010), and $^{238}\text{U}/^{235}\text{U} = 137.824$ mole ratio (Weyer et al., 2008), to calculate ^{234}U and ^{235}U respectively based on $[\text{U}]$.

Individual uncertainties for thorium were calculated to include contributions from (a) blank correction using the variance of the blanks measured over the course of the analyses, (b) standard error of the ratios of the analysis (typically close to counting statistics) and (c) spike calibration. For Pa we also included assessment of the correction from the yield correction, mass bias and machine background. In order to assess the reproducibility of the procedure, repeat analyses were performed on the GEOTRACES 2010-1 standard. For standards run with the GEOTRACES intercalibration and section samples the reproducibility for each isotope was 0.9% for ^{230}Th , 0.8% for ^{232}Th and 3.6% for ^{231}Pa .

Access restrictions:

None.

Primary references for this data set:

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